

Scalable Inference and Model Selection for Large-Scale Time Series with Complex Dependence Structures

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Introduction

Modern sensing technologies and digital systems generate **massive time series** with complex dependence structures.

In many applications, the goal is to perform **statistical inference**.

Classical likelihood-based methods do not scale well, which limits the analysis of long signals and/or large sensor networks.

Practitioners often rely on **approximations** or **ad hoc methods** that reduce computational cost but provide **limited or unclear statistical guarantees**.

Thesis Objective

Develop **scalable** statistical methods with **clear theoretical guarantees** for **reliable statistical inference** on **large dependent datasets**.

Applications include inertial sensor calibration in engineering and the estimation of tectonic rates using collections of position time series from large-scale GNSS station networks.

Chapter I: Accounting for Vibration Noise in Stochastic Measurement Errors of Inertial Sensors

This chapter corresponds to the following published article:

IEEE TRANSACTIONS ON SIGNAL PROCESSING, VOL. 72, 2024

2117

Accounting for Vibration Noise in Stochastic Measurement Errors of Inertial Sensors

Mucyo Karemera , Lionel Voirol , Davide A. Cucci , Wenfei Chu ,
Roberto Molinari , and Stéphane Guerrier 

Karemera, M.*, Voirol, L.*, Cucci, D. A., Chu, W., Molinari, R., Guerrier, S., "Accounting for Vibration Noise in Stochastic Measurement Errors of Inertial Sensors", *IEEE Transactions on Signal Processing*, 72, 2117–2129, 2024.

* indicates co-first authors and authors are listed alphabetically.

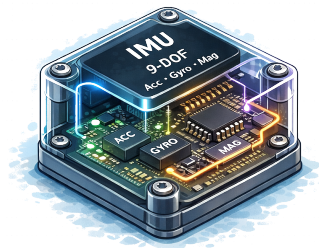
Inertial sensors are used in drones, smartphones and autonomous vehicles.

They measure acceleration and angular velocity at high frequency.

Their measurements are corrupted by errors that quickly accumulate, requiring the noise structure to be adequately modeled.

These signals and calibration data are then used in sensor-fusion algorithms, to estimate position, velocity, and attitude.

Measurement errors are generally modeled using **composite stochastic models**.



The GMWM estimator

Likelihood-based inference becomes **computationally prohibitive** for large and complex time series models.

The Generalized Method of Wavelet Moments¹ (GMWM) proposes to exploit the property of the **Wavelet Variance** (WV) to concisely capture the dependence structure of a time series in order to inverse the mapping between θ_0 and $\nu(\theta_0)$ to estimate θ_0 :

Let $\varepsilon_1, \dots, \varepsilon_n \sim F_{\theta_0}$ be a realization of the data-generating process with true parameter $\theta_0 \in \Theta$.

$$\hat{\theta} = \underset{\theta}{\operatorname{argmin}} \|\hat{\nu} - \nu(\theta)\|_{\Omega}^2,$$

where:

- $\hat{\nu}$ is the empirical wavelet variance² computed on $\varepsilon_1, \dots, \varepsilon_n$
- $\nu(\theta)$ is the model-implied wavelet variance
- Ω is a positive definite weighting matrix

¹ Guerrier, Stéphane et al., 2013, *Journal of the American Statistical Association*

² Percival, Donald P., 1995, *Biometrika*

The stochastic properties of device errors may vary with **external conditions**, such as **temperature** or **motion**³.

The impact of these dynamics on measurement errors is studied in **controlled environments**, such as rotation tables.

Vibration noise arises from internal mechanics, motion dynamics, or external conditions, and contaminates the measurement errors of sensing devices.

This type of perturbation is **not properly captured by standard stochastic error models**, leading to inaccurate calibration and degraded navigation performance.

³Stebler, Yannick et al., 2015, *IEEE Transactions on Instrumentation and Measurement*; Radi, Ahmed et al., 2018, *Gyroscopy and Navigation*

Chapter Objective

- ➊ Define a stochastic model that adequately captures vibration perturbation
- ➋ Derive its wavelet variance for GMWM-based estimation
- ➌ Establish the theoretical properties of the GMWM when incorporating this additional component

We consider a stochastic periodic process that we parameterize as follows:

$$S_t = \alpha \sin(\beta t + U) \quad (1)$$

with:

- amplitude $\alpha \in \mathbb{R}_+ = (0, +\infty)$,
- angular frequency $\beta \in (0, \pi]$ and
- phase $U \sim \mathcal{U}(0, 2\pi)$.

To make use of the GWMM, we need the theoretical WV of the processes defined in (1).

We derive the **theoretical WV** of S_t at scale j denoted as $\nu_j(\alpha, \beta)$:

$$\nu_j(\alpha, \beta) = \frac{\alpha^2 \left\{ 1 - \cos\left(\frac{\beta \tau_j}{2}\right) \right\}^2}{\tau_j^2 \{1 - \cos(\beta)\}}$$

where τ_j denotes the length of the wavelet filter at the j^{th} level of decomposition.

We consider a general class of **composite stochastic model**:

$$\varepsilon_t = W_t + Q_t + \sum_{k=1}^K Y_t^{(k)} + D_t + R_t + \sum_{l=1}^L S_t^{(l)},$$

where

- W_t is a White Noise,
- Q_t is a Quantization Noise,
- $Y_t^{(k)}$ is the k^{th} AR(1) process,
- D_t is a Deterministic Drift,
- R_t is a Random Walk and
- $S_t^{(l)}$ is the l^{th} sinusoidal process.

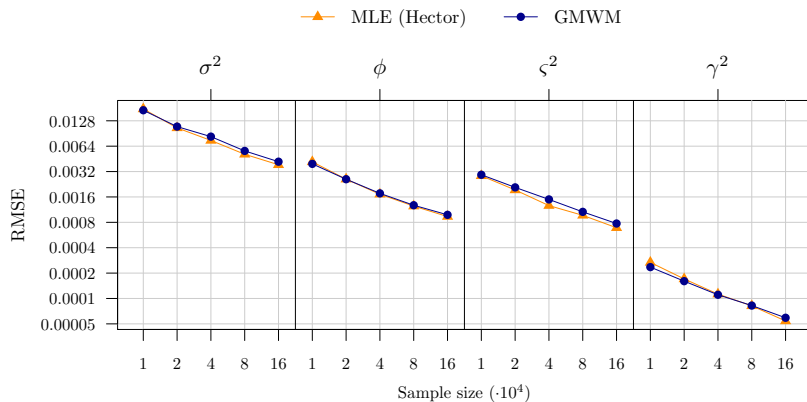


Figure: RMSE of the estimated parameters for the MLE (orange line) and the GMWM (blue line) for sample sizes $T = T_i \cdot 10^4$ with $\{T_1, \dots, T_5\} = \{1, 2, 4, 8, 16\}$

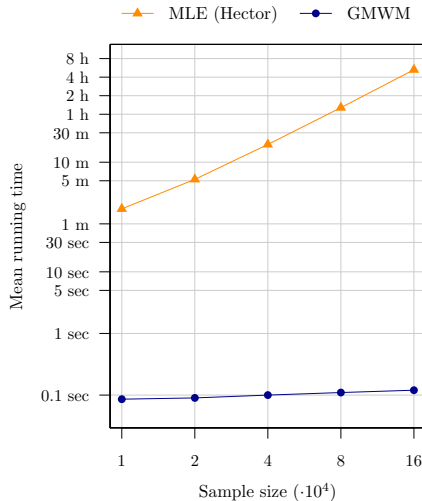
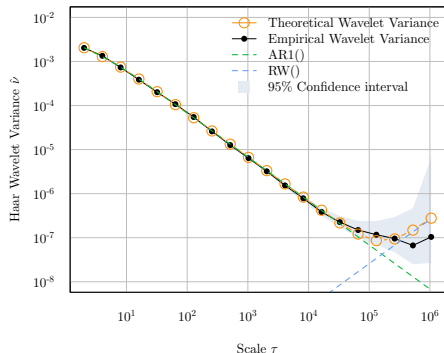


Figure: Mean running time of the MLE (orange line) and the GMWM (blue line) for sample sizes $T = T_i \cdot 10^4$ with $\{T_1, \dots, T_5\} = \{1, 2, 4, 8, 16\}$

We consider two signals collected from a GMAT IMU under two settings.

Fit at 0° s^{-1}



Fit at 200° s^{-1}

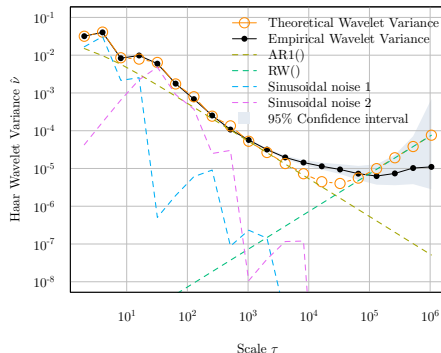


Figure: Empirical and model-implied wavelet variance of the estimated models for signals collected on a static IMU and when rotating at 200° s^{-1}

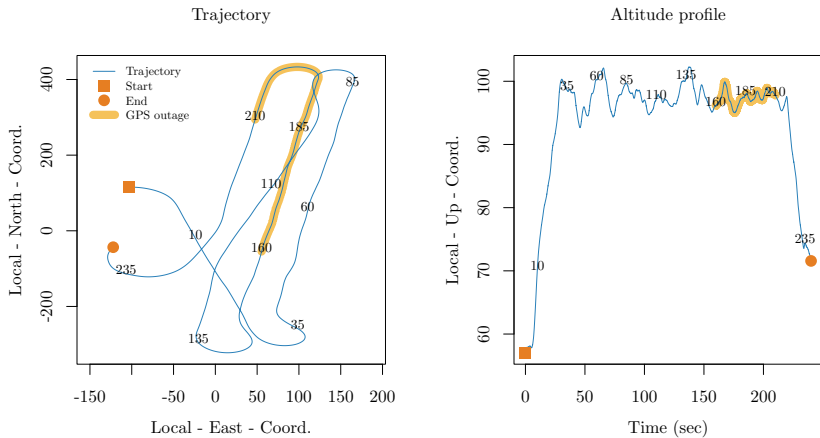


Figure: Trajectory in the emulation study

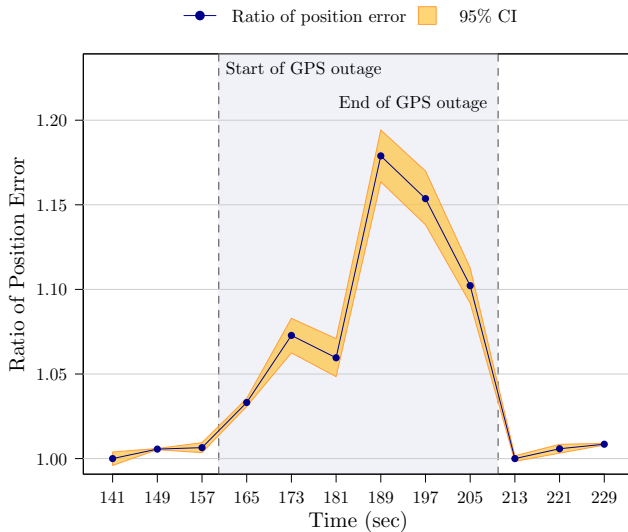


Figure: Mean ratio of position error

Chapter II: The Generalized Method of Wavelet Moments with eXogenous inputs

This chapter corresponds to the following published article:

Journal of Geodesy (2023) 97:14
<https://doi.org/10.1007/s00190-023-01702-8>

ORIGINAL ARTICLE



The Generalized Method of Wavelet Moments with eXogenous inputs: a fast approach for the analysis of GNSS position time series

Davide A. Cucci¹  · Lionel Voirol¹  · Gaël Kermarrec²  · Jean-Philippe Montillet^{3,4}  · Stéphane Guerrier^{1,5} 

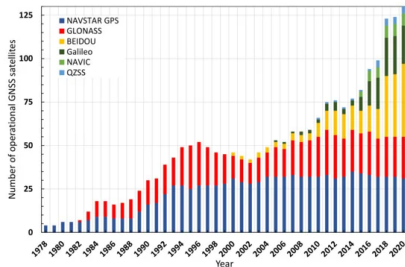
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Cucci, D. A.^{*}, Voirol, L.^{*}, Kermarrec, G., Montillet, J.-P., Guerrier, S., “*The Generalized Method of Wavelet Moments with eXogenous inputs: A Fast Approach for the Analysis of GNSS Position Time Series*”, *Journal of Geodesy*, 97(2), 2023.

^{*} indicates co-first authors and authors are listed alphabetically.

The **Global Navigation Satellite System (GNSS)** is an important tool to study multiple geophysical phenomena⁴ such as long-term tectonic movements⁵, fault strain accumulation, post-glacial rebound⁶, hydrological loading⁷ and crustal deformations⁸.

Networks of permanent stations record **sequences of position measurements** from various satellite constellations.



⁴Bock, Yehuda et al., 2016, *Reports on Progress in Physics*; Herring, T. A. et al., 2016, *Reviews of Geophysics*

⁵Bock, Yehuda et al., 2016, *Reports on Progress in Physics*

⁶Milne, Glenn A et al., 2001, *Science*

⁷Bevis, Michael et al., 2002, *Marine Geodesy*

⁸Williams, Simon, 2003, *Journal of Geophysical Research: Solid Earth*

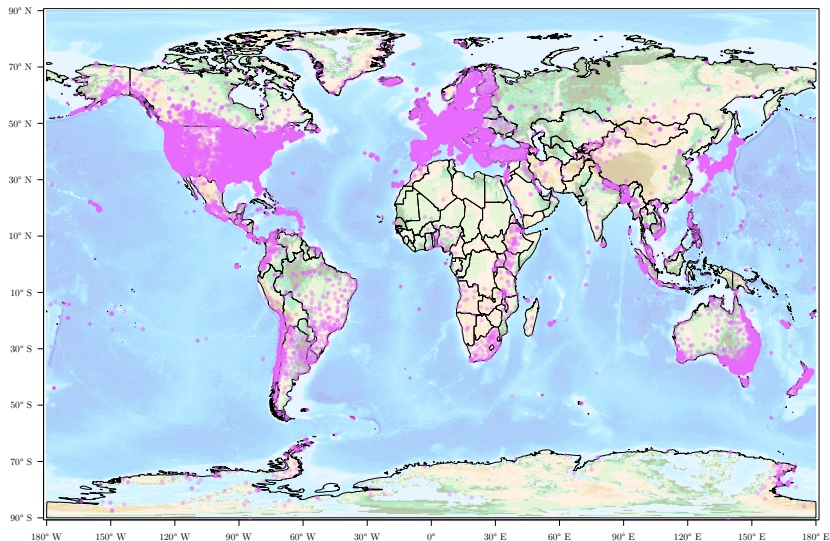


Figure: GNSS Stations processed by the Nevada Geodetic Laboratory

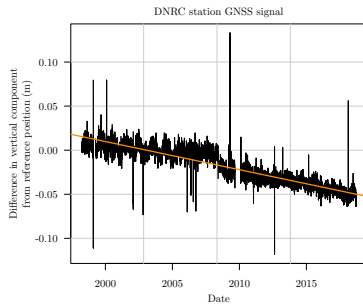
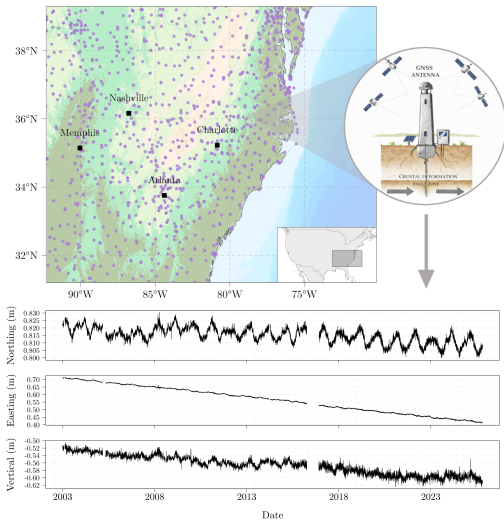


Figure: GNSS position time series and estimated trend

Figure: GNSS network along the U.S. east coast with example station and position time series

A general formulation of the models typically used to model GNSS signals can be expressed as

$$\mathbf{Y} = \mathbf{X}\boldsymbol{\beta}_0 + \boldsymbol{\varepsilon}, \text{ with } \boldsymbol{\varepsilon} \sim \mathcal{N}\{\mathbf{0}, \boldsymbol{\Sigma}(\boldsymbol{\gamma}_0)\}, \quad (2)$$

where

- $\mathbf{Y} \in \mathbb{R}^n$ is the response vector,
- $\mathbf{X} \in \mathbb{R}^{n \times p}$ is a fixed design matrix,
- $\boldsymbol{\beta}_0 \in \mathbb{R}^p$ is the vector of unknown parameters,
- $\boldsymbol{\varepsilon} \in \mathbb{R}^n$ is a zero-mean residual vector, with covariance matrix $\boldsymbol{\Sigma}(\boldsymbol{\gamma}_0)$, that depends on the unknown parameter vector $\boldsymbol{\gamma}_0 \in \mathbb{R}^q$. This matrix does not have a block diagonal structure and can be non-Toeplitz.

Hence, we define

$$\boldsymbol{\theta}_0 = \begin{bmatrix} \boldsymbol{\beta}_0^T & \boldsymbol{\gamma}_0^T \end{bmatrix}^T \in \boldsymbol{\Theta} \subset \mathbb{R}^{p+q}.$$

Trajectory Model:

With t_i representing the i^{th} time point, the mean function capturing trend, seasonal variations, offsets, and post-seismic relaxations is written as⁹:

$$\mathbb{E}[Y_i] = a + b(t_i - t_0) + \sum_{h=1}^m \left\{ c_h \sin(2\pi f_h t_i) + d_h \cos(2\pi f_h t_i) \right\} + \sum_{j=1}^r e_j H(t_i - t_j) + \sum_{k=1}^s l_k \left\{ 1 - \exp\left(\frac{-(t_i - t_k)}{\tau_k}\right) \right\} H(t_i - t_k), \quad (3)$$

for $i \in \{1, \dots, n\}$, and with $\beta_0 = \left[a, b, c_1, \dots, c_m, d_1, \dots, d_m, e_1, \dots, e_r, l_1, \dots, l_s \right]^T$.

Stochastic Model:

The stochastic component is modeled as a [combination of stochastic processes](#), typically coupling white noise with correlated processes such as Matérn or power-law processes.

⁹Langbein, John, 2008, *Journal of Geophysical Research: Solid Earth*; Bos, M. S. et al., 2008, *Journal of Geodesy*; Williams, Simon, 2003, *Journal of Geophysical Research: Solid Earth*

Maximum Likelihood Estimation: Computational Challenges

Inference is typically performed via **Maximum Likelihood Estimation (MLE)**.

The MLE is implemented in softwares such as CATS, Est_noise, and Hector¹⁰.

Each evaluation of the likelihood requires inversion of the $n \times n$ matrix $\Sigma(\gamma)$.

Computational complexity: $\mathcal{O}(n^\delta)$, $\delta \in [2, 3]$.

In practice, the analysis of complex geodynamic processes requires to estimate signals from thousands of GNSS stations which record daily observations over decades and where different noise models must be tested¹¹.

This procedure which has to be performed routinely becomes **impractical** due to the **large amount (e.g., weeks) of processing time** required¹².

¹⁰Williams, Simon, 2008, *GPS Solutions*; Langbein, John, 2008, *Journal of Geophysical Research: Solid Earth*; Bos, M. S. et al., 2008, *Journal of Geodesy*

¹¹He, Xiaoxing et al., 2021, *Remote Sensing*

¹²He, X et al., 2019, *Journal of Geodesy*; Bos, Machiel S. et al., 2020; Tunini, Lavinia et al., 2024, *Earth System Science Data*

Chapter Objective

Develop a **computationally efficient** alternative to the MLE for **linear models with dependent errors**, and apply it to large collections of **GNSS position time series** for estimating horizontal tectonic velocities and vertical crustal uplift rates.

We propose a **two-step estimation procedure** combining a **Generalized Least Squares (GLS)** approach and the **GMWM** that we introduce as **GMWMX** for Generalized Method of Wavelet Moments with eXogenous variables.

Step 1: Estimation of β_0 (GLS step)

Given a covariance matrix Σ , define

$$\hat{\beta} = \left(\mathbf{X}^\top \Sigma^{-1} \mathbf{X} \right)^{-1} \mathbf{X}^\top \Sigma^{-1} \mathbf{Y},$$

and the estimated residuals computed using $\hat{\beta}$:

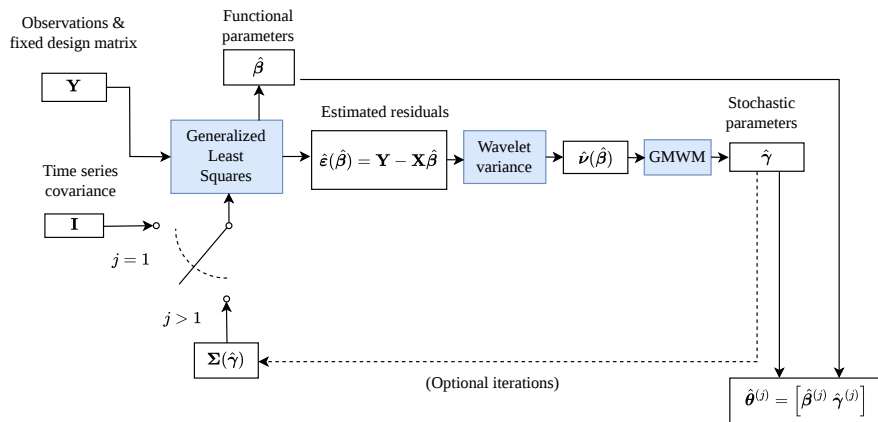
$$\hat{\varepsilon}(\hat{\beta}) = \mathbf{Y} - \mathbf{X}\hat{\beta}.$$

Step 2: Estimation of γ_0 (GMWM step)

Consider $\hat{\nu}(\hat{\beta})$, the empirical WV computed on $\hat{\varepsilon}(\hat{\beta})$ and, $\nu(\gamma)$, the model-implied WV. We define:

$$\hat{\gamma} = \underset{\gamma}{\operatorname{argmin}} \left\{ \hat{\nu}(\hat{\beta}) - \nu(\gamma) \right\}^\top \Omega \left\{ \hat{\nu}(\hat{\beta}) - \nu(\gamma) \right\}.$$

GMWMX: Iterative algorithm



We denote the estimator with one or two iterations as the GMWMX-1 and GMWMX-2, respectively.

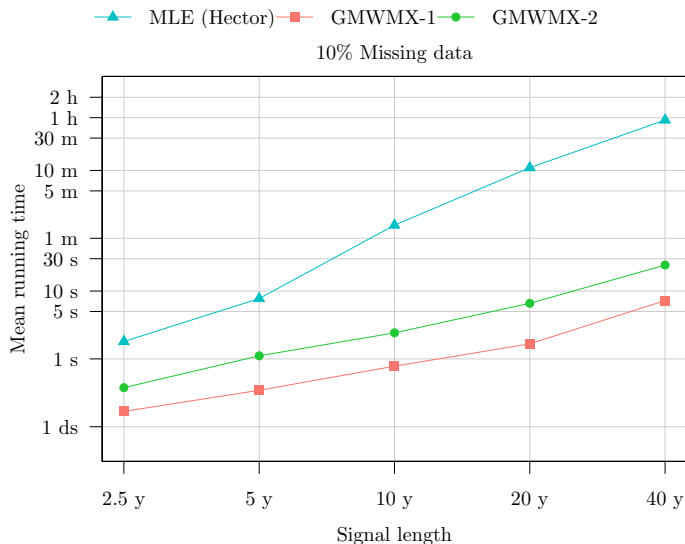


Figure: Mean running time of the MLE, the GMWMX-1 and the GMWMX-2 as a function of the sample size.

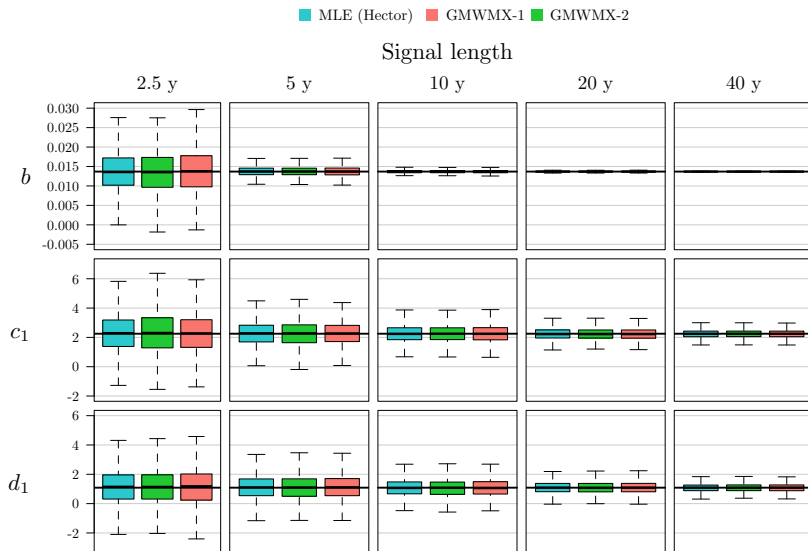


Figure: Boxplots of the estimated parameters of the model with the GMWMX-1, the GMWMX-2 and the MLE

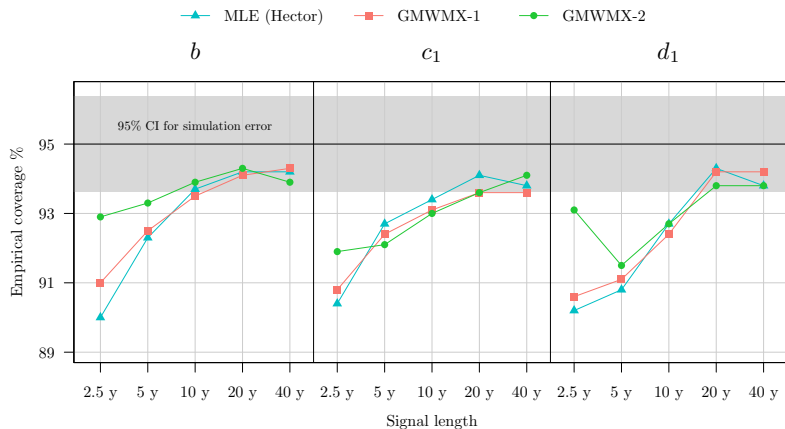


Figure: Empirical coverage of the confidence intervals at level $1 - \alpha = 0.95$ for the parameters b , c_1 and d_1 for GMWMX-1, GMWMX-2 and the MLE as a function of sample size. The grey area represents a 95% confidence interval of the simulation error.

Case Study

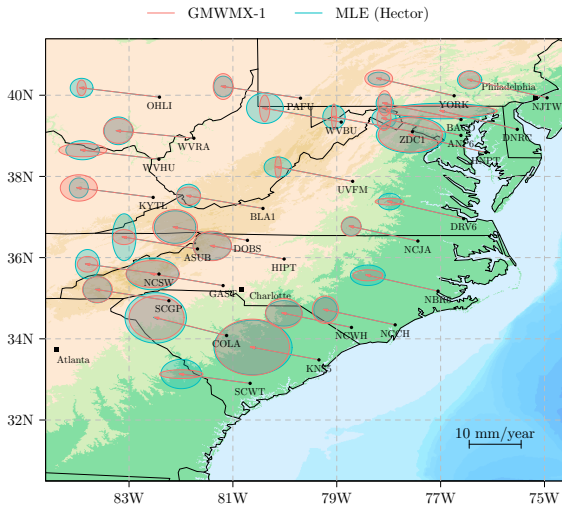


Figure: Estimated North-East velocity solutions for 33 GNSS receivers distributed over the East coast of the USA using the GMWMX-1 and Hector software (MLE).

This chapter is based on the following manuscript:

Towards Open Science: Monitoring Crustal Deformations in North America

Lionel Voirol¹, Haotian Xu², Yuming Zhang³, Luca Insolia^{4,5},
Roberto Molinari² & Stéphane Guerrier^{4,5}

¹Geneva School of Economics and Management, University of Geneva, Switzerland

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⁴School of Pharmaceutical Sciences, University of Geneva, Switzerland

⁵Department of Earth Sciences, University of Geneva, Switzerland

Voirol, L., Xu, H., Zhang, Y., Insolia, L., Molinari, R., Guerrier, S., “Towards Open Science: Monitoring Crustal Deformations in North America”.

GNSS position time series often contain **missing observations** because of receiver failures, maintenance, or poor observation conditions¹³.

Valid inference depends on the **choice of the stochastic model**¹⁴. Misspecified noise structures can produce **incorrect standard errors** by factors ranging from 5 to 11¹⁵.

Model selection becomes part of the inferential problem.

In practice, this is usually done by fitting several candidate stochastic models and selecting one with likelihood-based criteria such as AIC or BIC¹⁶.

This adds a substantial **computational burden** to large-scale GNSS analyses.

¹³Bos, M. S. et al., 2008, *Journal of Geodesy*; Bos, MS et al., 2013, *Journal of Geodesy*

¹⁴Langbein, John et al., 1997, *Journal of Geophysical Research: Solid Earth*; Williams, Simon, 2003, *Journal of Geodesy*

¹⁵Mao, Ailin et al., 1999, *Journal of Geophysical Research: Solid Earth*; Wang, Wei et al., 2012, *Advances in space research*

¹⁶He, Xiaoxing et al., 2017, *Journal of Geodynamics*; He, X et al., 2019, *Journal of Geodesy*

Chapter Objective

- ① Incorporate a **missingness mechanism** directly into the modeling framework
- ② Develop a **computationally efficient model selection criterion** for the stochastic noise structure
- ③ Study in depth the **theoretical properties** of the resulting estimator

Recall the model previously defined as:

$$\mathbf{Y} = \mathbf{X}\boldsymbol{\beta}_0 + \boldsymbol{\varepsilon}, \text{ with } \boldsymbol{\varepsilon} \sim \mathcal{N}\{\mathbf{0}, \boldsymbol{\Sigma}(\boldsymbol{\gamma}_0)\}.$$

A reasonable strategy to model missing data is to consider the variable $\tilde{\mathbf{Y}}$ related to the latent variable $\mathbf{Y}$ by the following mechanism:

$$\tilde{\mathbf{Y}} = \mathbf{Z} \odot \mathbf{Y},$$

where $\odot$ denotes the Hadamard product and $\mathbf{Z} = [Z_1, \dots, Z_n]^T$ is a binary-valued stationary process independent of $\mathbf{Y}$ which encodes observed entries by 1 and 0 otherwise, with $\mathbb{E}[Z_i] = \mu(\boldsymbol{\vartheta}_0) \in (0, 1]$ and covariance matrix $\boldsymbol{\Lambda}(\boldsymbol{\vartheta}_0) \in \mathbb{R}^{n \times n}$, parametrized by $\boldsymbol{\vartheta}_0 \in \boldsymbol{\Upsilon} \subset \mathbb{R}^k$.

We consider a general class of models for $\mathbf{Z}$ as:

$$Z_i = \begin{cases} Z_{i-1}, & \text{with probability } \rho, \\ W_i \sim \text{Bernoulli}(\mu(\boldsymbol{\vartheta}_0)), & \text{with probability } 1 - \rho, \end{cases} \quad (4)$$

With this new parametrization, we define

$$\boldsymbol{\theta}_0 = \left[\boldsymbol{\beta}_0^{\text{T}}, \boldsymbol{\gamma}_0^{\text{T}}, \boldsymbol{\vartheta}_0^{\text{T}} \right]^{\text{T}}$$

and $\tilde{\mathbf{X}} = \left(\mathbf{Z} \otimes \mathbf{1}_p^{\text{T}} \right) \odot \mathbf{X} \in \mathbb{R}^{n \times p}$ as the design matrix $\mathbf{X}$ with zero-valued rows for missing observations in $\mathbf{Y}$, where $\otimes$ denotes the Kronecker product and $\mathbf{1}_p$ is a vector of ones of dimension p .

We propose a three-steps statistical procedure that we introduce as **WAMORE** for Wavelet Moment Regression.

Step 1: Estimation of ϑ_0

Using $\mathbf{Z}$, we first estimate ϑ_0 by Maximum Likelihood. Let $\hat{\vartheta}$ denote our estimator.

Step 2: Estimation of β_0

We then consider a Least Squares approach to obtain an estimate of β_0 and define:

$$\hat{\beta} = \left(\tilde{\mathbf{X}}^T \tilde{\mathbf{X}} \right)^{-1} \tilde{\mathbf{X}}^T \tilde{\mathbf{Y}}.$$

Step 3: Estimation of γ_0

Let $\hat{\nu}$ represent the empirical WV computed on the estimated residuals process $\hat{\varepsilon} = \tilde{\mathbf{Y}} - \tilde{\mathbf{X}}\hat{\beta}$ and $\nu(\gamma, \vartheta)$ the theoretical WV of $\tilde{\varepsilon} = \tilde{\mathbf{Y}} - \tilde{\mathbf{X}}\beta = \varepsilon \odot \mathbf{Z}$. Given $\hat{\vartheta}$, we estimate γ_0 with:

$$\hat{\gamma} = \underset{\gamma}{\operatorname{argmin}} \left\| \hat{\nu} - \nu(\gamma, \hat{\vartheta}) \right\|_{\Omega}^2 \quad (5)$$

We aim to construct a computationally efficient model selection criterion for the **stochastic model**.

Recall the error term $\varepsilon \sim \mathcal{N}\{\mathbf{0}, \mathbf{\Sigma}(\gamma_0)\}$ with $\gamma_0 \subset \mathbb{R}^q$ and consider $\hat{\nu}$, the empirical WV computed on ε .

A reasonable fit metric is given by:

$$C = \mathbb{E}_0 \left[\mathbb{E} \left\{ \|\hat{\nu}_0 - \nu(\hat{\gamma})\|_{\mathbf{\Omega}}^2 \right\} \right] = \mathbb{E}_0 \left(\mathbb{E} \left[\{\hat{\nu}_0 - \nu(\hat{\gamma})\}^T \mathbf{\Omega} \{\hat{\nu}_0 - \nu(\hat{\gamma})\} \right] \right),$$

where $\hat{\gamma}$ is computed on the sample ε and $\hat{\nu}_0$ is computed on an independent copy, ε_0 satisfying $\varepsilon \stackrel{d}{=} \varepsilon_0$.

It can be shown that:

$$C = \mathbb{E} \left[\|\hat{\boldsymbol{\nu}} - \boldsymbol{\nu}(\hat{\boldsymbol{\gamma}})\|_{\boldsymbol{\Omega}}^2 \right] + 2\text{tr} \left[\text{cov} \{ \hat{\boldsymbol{\nu}}, \boldsymbol{\nu}(\hat{\boldsymbol{\gamma}}) \} \boldsymbol{\Omega} \right]$$

Using a first-order approximation of $\boldsymbol{\nu}(\hat{\boldsymbol{\gamma}})$ around $\boldsymbol{\gamma}_0$, an estimator of C is obtained. With $\boldsymbol{\Omega} = \mathbf{V}^{-1}$, the resulting complexity penalty simplifies to:

$$\hat{C} = \|\hat{\boldsymbol{\nu}} - \boldsymbol{\nu}(\hat{\boldsymbol{\gamma}})\|_{\boldsymbol{\Omega}}^2 + 2q$$

The empirical WV as a quadratic form

In order to obtain $\nu(\gamma, \vartheta)$ to estimate γ_0 in (5) and $\mathbf{V} = \text{var}(\hat{\nu})$ to compute $\hat{C}$, we need to derive their explicit form.

Denoting the vector of Wavelet coefficients at scale j as $\mathbf{W}_j$ computed on a signal $\varepsilon \in \mathbb{R}^n$ with covariance matrix $\Sigma(\gamma_0)$, we can write:

$$\mathbf{W}_j = \begin{bmatrix} W_{j,1} \\ W_{j,2} \\ \vdots \\ W_{j,M_j} \end{bmatrix} = \mathbf{B}_j \varepsilon = \begin{bmatrix} \mathbf{h}_j^T & 0 & \dots & 0 \\ 0 & \mathbf{h}_j^T & \dots & 0 \\ \vdots & \vdots & \ddots & \vdots \\ 0 & 0 & \dots & \mathbf{h}_j^T \end{bmatrix} \begin{bmatrix} \varepsilon_1 \\ \varepsilon_2 \\ \vdots \\ \varepsilon_n \end{bmatrix}$$

where $\mathbf{B}_j$ is an $M_j \times n$ matrix and $\mathbf{h}_j^T$ is a vector of length 2^j containing the wavelet filter. Hence, we can rewrite the empirical WV as a quadratic form:

$$\hat{\nu}_j = \frac{1}{M_j} \sum_{i=1}^{M_j} W_{j,i}^2 = \frac{1}{M_j} \mathbf{W}_j^T \mathbf{W}_j = \frac{1}{M_j} (\mathbf{B}_j \varepsilon)^T \mathbf{B}_j \varepsilon = \varepsilon^T \mathbf{A}_j \varepsilon, \text{ where } \mathbf{A}_j = \frac{\mathbf{B}_j^T \mathbf{B}_j}{M_j}$$

Using the quadratic form representation, we obtain the **theoretical WV**:

$$\mathbb{E}(\hat{\nu}_j) = \nu_j = \mathbb{E} \left[\boldsymbol{\varepsilon}^T \mathbf{A}_j \boldsymbol{\varepsilon} \right] = \text{tr} \{ \mathbf{A}_j \boldsymbol{\Sigma}(\gamma_0) \},$$

Using the Gaussian quadratic-form result, we obtain $\mathbf{V}$, the **covariance of the empirical WV $\hat{\nu}$** :

$$\text{cov}\{\hat{\nu}_i, \hat{\nu}_j\} = \text{cov}(\boldsymbol{\varepsilon}^T \mathbf{A}_i \boldsymbol{\varepsilon}, \boldsymbol{\varepsilon}^T \mathbf{A}_j \boldsymbol{\varepsilon}) = 2\text{tr}\{\mathbf{A}_i \boldsymbol{\Sigma}(\gamma_0) \mathbf{A}_j \boldsymbol{\Sigma}(\gamma_0)\}.$$

First explicit form of $\mathbf{V}$ ¹⁷.

We provide a **computationally efficient algorithm** to obtain $\text{cov}\{\hat{\nu}_i, \hat{\nu}_j\}$ allowing to compute these quantities without passing by their matrix expression.

¹⁷Zhang, Nien Fan, 2008, *Metrologia*

Inference on the functional parameters

- Consistent estimators of β_0 and γ_0
- Asymptotic normality of $\hat{\beta}$ under short- and long-memory dependence regime
- Valid confidence intervals under dependent errors and missing observations

Model selection

- Consistent estimator of the target criterion: $\hat{C} = C + o_p(1)$
- Same asymptotic overfitting and underfitting probabilities as the AIC

Simulation Studies

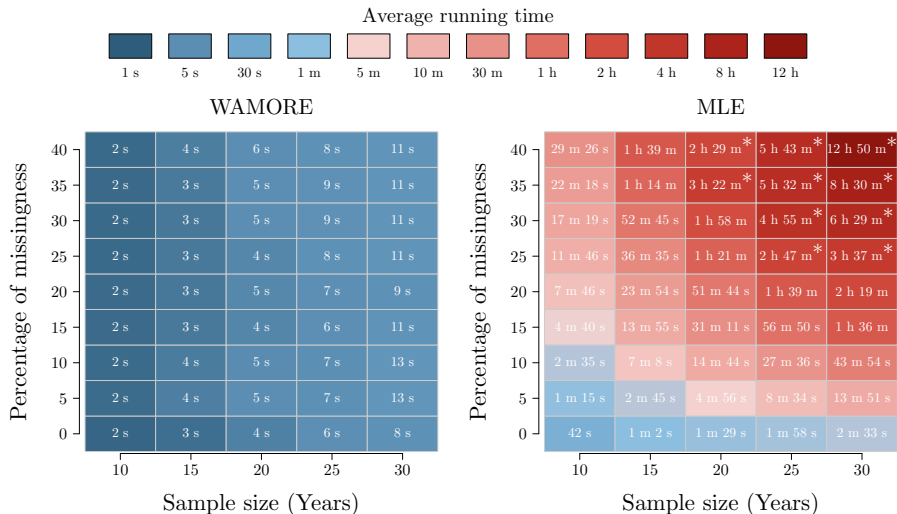


Figure: Average running time of the WAMORE (left panel) vs the MLE (right panel) for the setting with the WN + MAT + RW model. Settings annotated with a “*” were evaluated only on 5 Monte Carlo realizations due to computational constraints.

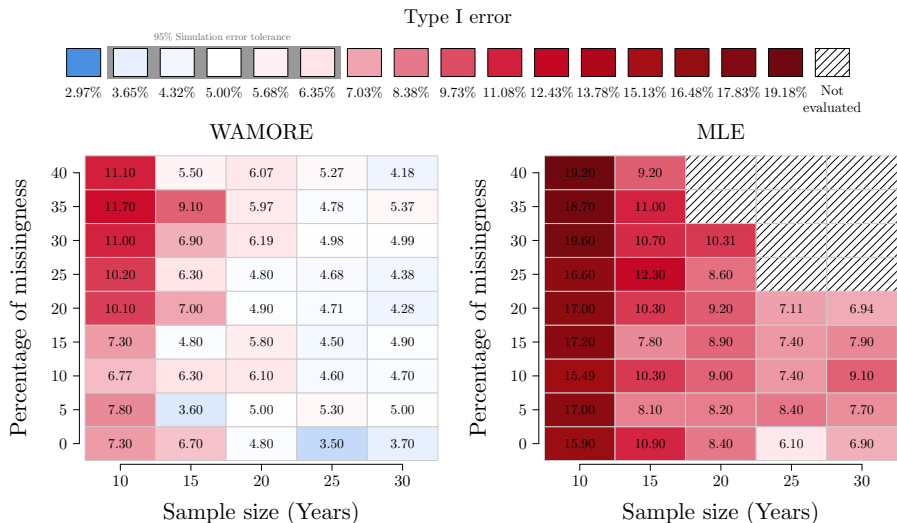


Figure: Empirical type I error (in %) of the WAMORE (left panel) vs the MLE (right panel) for the setting with the WN + MAT + RW model.

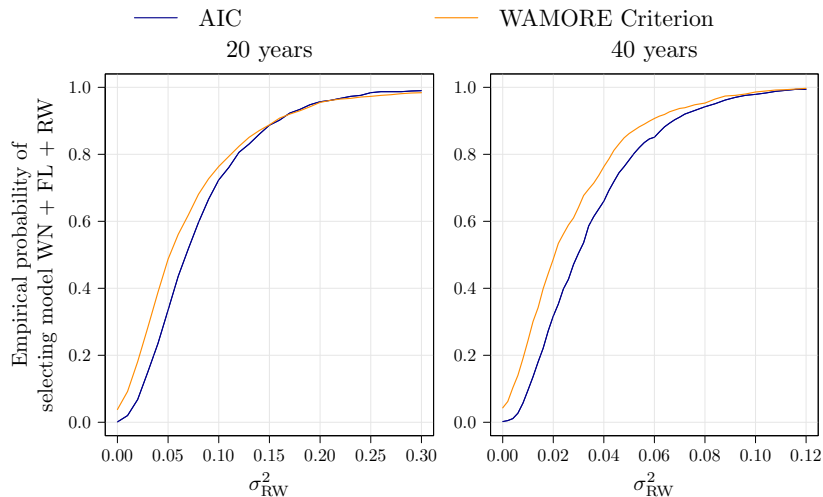


Figure: Selection rate of the AIC and the proposed WAMORE criterion for setting considering 20 years of daily data (left panel) and 40 years of daily data (right panel).

Case Study

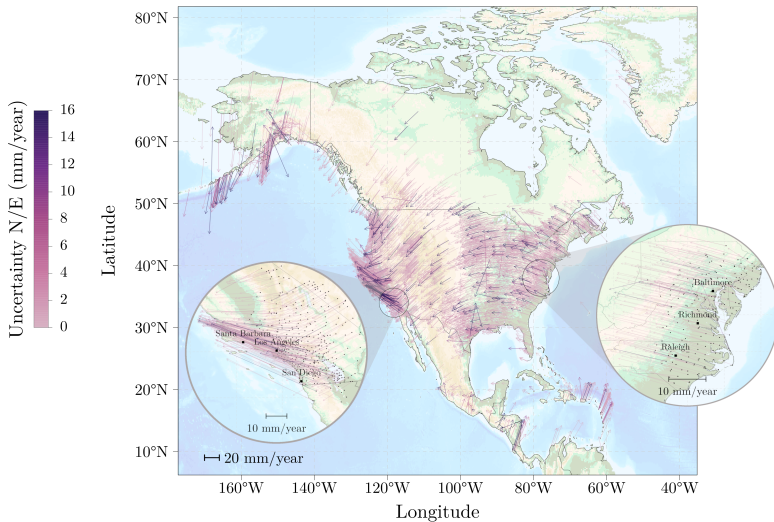


Figure: Estimated tectonic velocities and associated uncertainty using the WAMORE methodology. Arrow length encodes velocity magnitude, and color represents uncertainty, computed as the ℓ_2 -norm of the north and east component standard errors.

Case Study

<https://data-analytics-lab.shinyapps.io/wamore-crustal-deformation-explorer>



Figure: Interactive WAMORE crustal deformations explorer.

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Thank You!



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